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Atmospheric Chemistry and Physics Discussions

# Nitrous oxide emissions from the Arabian Sea: A synthesis

**H. W. Bange[1,7], M. O. Andreae[1], S. Lal[2], C. S. Law[3], S. W. A. Naqvi[4], P. K. Patra[2,8], T. Rixen[5,9], and R. C. Upstill-Goddard[6]**

[1]Max Planck Institute for Chemistry, Mainz, Germany
[2]Physical Research Laboratory, Ahmedabad, India
[3]Plymouth Marine Laboratory, Plymouth, United Kingdom
[4]National Institute of Oceanography, Dona Paula, Goa, India
[5]University of Hamburg, Hamburg, Germany
[6]University of Newcastle, Newcastle upon Tyne, United Kingdom
[7]now at Institute for Marine Research, Kiel, Germany
[8]now at Frontier Research System for Global Change, Yokohama, Japan
[9]now at Center for Tropical Marine Ecology, Bremen, Germany

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Correspondence to: H. W. Bange (hbange@ifm.uni-kiel.de)




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## Abstract

We computed high-resolution (1° latitude × 1° longitude) seasonal and annual nitrous oxide ($N_2O$) concentration fields for the Arabian Sea surface layer using a database containing more than 2400 values measured between December 1977 and July 1997. $N_2O$ concentrations are highest during the southwest (SW) monsoon along the southern Indian continental shelf. Annual emissions range from 0.33 to 0.70 Tg $N_2O$ and are dominated by fluxes from coastal regions during the SW and northeast monsoons. However, the tendency to focus on measurements in locally restricted features in combination with insufficient seasonal data coverage leads to considerable uncertainties of the concentration fields and thus in the flux estimates, especially in the coastal zones of the northern and eastern Arabian Sea.

## 1. Introduction

Nitrous oxide ($N_2O$) is an atmospheric trace gas that influences, directly and indirectly, the Earth's climate (Prather et al., 2001). Source estimates indicate that the world's oceans play a major role in the global budget of atmospheric $N_2O$ (Seitzinger et al., 2000). Upwelling regions, such as the eastern tropical Pacific and the Arabian Sea, are sites of high $N_2O$ production via denitrification and/or nitrification processes that occur at the boundaries of the oxygen depleted water masses (Codispoti et al., 1992). Following the studies of Law and Owens (1990) and Naqvi and Noronha (1991), it has been speculated that the Arabian Sea, especially its upwelling-dominated northwestern part, represents a hot spot for $N_2O$ emissions and makes a substantial contribution to the global budget of atmospheric $N_2O$. However, the situation is apparently somewhat more complicated, because recent data show seasonal $N_2O$ emissions from the continental shelf area of India also to be important (Naqvi et al., 2000). Previous $N_2O$ flux estimates are compromised by significant temporal and spatial biases. Moreover, we recognize that in efforts to model global oceanic $N_2O$ emissions, the Arabian



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Sea appears to be under-represented mainly owing to the relatively low spatial resolution of the applied models and/or missing data from this region (Nevison et al., 1995; Seitzinger et al., 2000; Suntharalingam and Sarmiento, 2000). Here we present a comprehensive compilation of $N_2O$ measurements from the Arabian Sea surface layer from 1977 to 1997. These data were used to calculate mean seasonal and annual climatological $N_2O$ fields with a 1° latitude × 1° longitude resolution. On the basis of the $N_2O$ surface concentration fields, $N_2O$ emissions from the Arabian Sea were reassessed.

## 2. Data sources

For our study we compiled $N_2O$ measurements from 0–10 m water depth within the study area (44°−80° E, 0°−27° N) excluding the Persian Gulf and the Red Sea (Fig. 1). The majority of the data were collected during individual national contributions to the international Joint Global Ocean Flux Study (JGOFS) – Arabian Sea Process Study between 1994 and 1997. Pre-JGOFS data were from cruises in 1977/1978, 1986, and 1988. An overview of the data sources is given in Table 1. (Unfortunately, data from the 1992 Netherlands Indian Ocean Program were unavailable for this reassessment.) $N_2O$ concentrations are typically reported in nmol $L^{-1}$, however, the data listed in Weiss et al. (1992) are in dry mole fractions. We recalculated the Weiss et al. (1992) $N_2O$ concentrations with the reported water temperature, a mean seasonal salinity of 35.75, as calculated from climatological salinity data (see below), and an atmospheric pressure of 1 atm (Weiss and Price, 1980). We are aware that this procedure introduces an additional error; however, the dependence of the $N_2O$ solubility on salinity and pressure is small and the resulting uncertainty of about ± 1% is acceptable for our purposes.

Weekly averaged wind speeds for the period July 1987 to December 1995 were derived from satellite-based Special Sensor Microwave / Imager measurements by using an algorithm developed by Schlüssel (1995). Weekly composites of 18 km × 18 km gridded, day and night multichannel sea surface temperatures (SSTs) satellite data




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for the period 1986 to 1995 were provided by the Physical Oceanography Distributed Active Archive Center of the Jet Propulsion Laboratory, California Institute of Technology, Pasadena, California (http://podaac.jpl.nasa.gov:2031/DATASET_DOCS/avhrr_wkly_mcsst.html). Monthly climatological salinities with a resolution of 1°× 1° were obtained from the World Ocean Atlas 1998 (http://www.nodc.noaa.gov/OC5/data_woa.html).

## 3. $N_2O$ fields

For the calculation of the $N_2O$ fields we applied a modified procedure originally described by Conkright et al. (1994) and further developed by Kettle et al. (1999). The original data sets were combined to form a database with 2463 values. The database was then divided into 12 monthly databases. A statistical checking procedure was implemented, wherein the monthly database values were pooled into 5°× 5° squares. For each 5°× 5° square a mean and standard deviation (sd) were calculated and individual data were compared with the mean. Values falling outside 3 times the sd of the mean were omitted and the procedure was repeated until no further values were eliminated. In squares with 3 values or fewer, the checking procedure was omitted and the remaining values accepted. This procedure removed 49 data points. The modified monthly databases were then subdivided into 1°× 1° squares. Mean $N_2O$ values (so-called $N_2O$ pixels) were calculated from the data in each square. If there was only one value within the square, it was accepted as a pixel. Monthly $N_2O$ pixel data sets were then combined into four seasonal sets: northeast (NE) monsoon (December to February, DJF), intermonsoon (March to May, MAM), southwest (SW) monsoon (June to August, JJA), and intermonsoon (September to November, SON) (Figs. 2a–d). Finally, the four seasonal sets were combined to form an annual $N_2O$ pixel set. For the annual and for each of the four seasonal and pixel sets, we calculated means for Arabian Sea biogeographic provinces, i.e. the Northwestern Arabian Upwelling, Indian Monsoon Gyres, and Western India Coastal provinces (INDW) (Longhurst, 1998).




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The biogeographic means were used to create a 1°× 1° first-guess field which was smoothed with a 9-point 2-dimensional operator (Shuman, 1957). A 1°× 1° correction field was computed for each of the seasonal and annual $N_2O$ pixel data by applying the distance-weighted interpolation scheme of Conkright et al. (1994). In order to preclude any smoothing of small-scale features, we reduced the influence radius from 555 km to 222 km. The correction field was then added to the first-guess field and smoothed (Shuman, 1957), yielding the final 1°× 1° annual and seasonal $N_2O$ fields. The final seasonal and annual $N_2O$ concentration fields are available from the German JGOFS data management (http://www.ifm.uni-kiel.de/jgofs/dm).

## 4. Air-sea exchange

The air-sea exchange flux density ($F$) was parameterized as

$$F = k_w(u)(C_w - C_a),$$

where $k_w$ is the gas transfer coefficient as a function of wind speed ($u$), $C_w$ is the $N_2O$ seawater concentration, and $C_a$ is the equilibrium $N_2O$ concentration in seawater. $C_a$ was calculated using

$$C_w + \beta(T,S)x'P,$$

where $x'$ is the atmospheric $N_2O$ dry mole fraction, $P$ is the atmospheric pressure, and $\beta$ is the Bunsen solubility, which is a function of the water temperature ($T$) and salinity ($S$) (Weiss and Price, 1980). To calculate $k_w$, we used the tri-linear $k_w - u$ relationship of Liss and Merlivat (1986) (LM86), the quadratic $k_w - u$ relationship for climatological wind data of Wanninkhof (1992) (W92), and the combined linear and quadratic $k_w - u$ relationship from Nightingale et al. (2000) (N00). $k_w$ was adjusted by multiplying with $(Sc/600)^{-n}$ ($n$ = 2/3 for wind speeds <3.6 m $s^{-1}$ and $n$ = 1/2 for wind speeds >3.6 m $s^{-1}$) for LM86, $(Sc/660)^{-0.5}$ for W92, and $(Sc/660)^{-0.5}$ for N00, where $Sc$ is the Schmidt number for $N_2O$. Sc was calculated using empirical equations for the



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kinematic viscosity of seawater (Siedler and Peters, 1986) and the diffusion coefficient of $N_2O$ in water. The $N_2O$ diffusion coefficients ($\boldsymbol{D}_{N2O}$ in $m^2\ s^{-1}$) were calculated with Eq. (1) derived from the data given in Broecker and Peng (1974) and, alternatively, with the new Eq. (2) derived from a compilation of actual measurements (Rhee, 2000):

$$\log_{10} \boldsymbol{D}_{N2O} = -1008.28/RT - 5.245 \quad (1)$$

$$\boldsymbol{D}_{N2O} = 3.16 \times 10^{-6} \exp(-18370/RT), \quad (2)$$

where $T$ is the water temperature in $K$ and $R$ is the universal gas constant.

$C_w$ was taken from the 1°× 1° seasonal $N_2O$ fields (DJF, MAM, JJA, SON). For the calculation of $\beta$, $Sc$, and $k_w$, seasonal 1°× 1° fields of wind speed, SST, and salinity were computed from the data sources given above. Atmospheric pressure was set to 1 atm. A mean $x'$ of 307 ppb for the period July 1978–July 1997 was calculated from the monthly mean values observed at the Cape Grim (Tasmania) and Adrigole/Mace Head (Ireland) monitoring stations of the ALE/GAGE/AGAGE program (updated version July 2000). The data are available from the anonymous ftp site cdiac.esd.ornl.edu (subdirectory /pub/ale_gage_Agage/Agage/monthly) at the Carbon Dioxide Information Analysis Center in Oak Ridge, Tennessee. $N_2O$ fluxes were calculated by multiplying the area of a 1°× 1° square with its flux density calculated as described above. The sum of the $N_2O$ fluxes of the 1°× 1° squares yields the total $N_2O$ emissions from the Arabian Sea (surface area: $6.8 \times 10^{12}$ $m^2$).

## 5. Results and discussion

Derived seasonal $N_2O$ concentration fields are shown in Fig. 3. Elevated $N_2O$ concentrations occur in coastal areas of the Arabian Sea during JJA (Fig. 3c). During DJF, $N_2O$ is higher in the eastern than in the western Arabian Sea, whereas during MAM and SON these $N_2O$ distributions are rather similar (Figs. 3b and 3c). However, the SON database is comparatively small, lending a note of caution to such a conclusion (Fig. 2d). The seasonal variability in $N_2O$ concentrations is clearly dominated




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by coastal upwelling in the Arabian Sea. During the SW monsoon, $N_2O$-rich subsurface waters are brought to the surface layer (see e.g., Bange et al., 2000; Patra et al., 1999). Interestingly, maximum $N_2O$ concentrations are found on the eastern Indian continental shelf, consistent with the observations by Patra et al. (1999). However, the calculated $N_2O$ values in the eastern Arabian Sea ($>70^\circ$ E) during JJA and SON, 5.8 to 36.3 nmol $L^{-1}$ (Figs. 3c and 3d), are considerably lower than the 5.3–436 nmol $L^{-1}$ range recently reported by Naqvi et al. (2000). It is possible that the enormous $N_2O$ accumulation observed along the Indian coast during the late summer and autumn is in part due to an (anthropogenic?) intensification of the natural coastal hypoxic system as a shift to anoxic conditions in the subsurface layers appears to have occurred in recent years (Naqvi et al., 2000). But if the $N_2O$ concentrations were high even before this intensification, then our analysis would underestimate the $N_2O$ concentrations and the associated fluxes from this region, especially during SON (see below).

Annual $N_2O$ emissions computed as the sum of the seasonal $N_2O$ emissions range from 0.37 to 0.78 Tg $N_2O$ $yr^{-1}$, depending on which air-sea transfer parameterization is used (Table 2). The use of the $N_2O$ diffusion coefficient of Rhee (2000) yielded about 10% lower $N_2O$ emissions (Table 3). Thus, we conclude that previous estimates using the $N_2O$ diffusion coefficient of Broecker and Peng (1974) may be overestimated.

$N_2O$ emissions during the SW monsoon (JJA) dominate the annual emissions, accounting for about 64–70% of the total. The second largest contribution occurs during the NE monsoon (DJF) (21–26%), whereas emissions from the intermonsoon period MAM seems to be of minor importance (2–3%). Our revised estimate for the annual $N_2O$ flux from the Arabian Sea is much more tightly constrained than the previous consensus of 0.16–1.5 Tg $N_2O$ $yr^{-1}$ derived using averaged in-situ data from a smaller number of studies (Table 4) (Bange et al., 1996a; Bange et al., 2000; Lal and Patra, 1998; Law and Owens, 1990; Naqvi and Noronha, 1991; Upstill-Goddard et al., 1999). The data listed in Table 4 depict the "historical" development of published $N_2O$ flux estimates for the Arabian Sea and show a considerable divergence. However, the fluxes listed are difficult to compare since they were extrapolated to different Arabian Sea sur-




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face areas and partly biased by the use of non-seasonal data sets and limited spatial data coverage.

## 6. Conclusions

Our calculated seasonal $N_2O$ concentration fields and associated air-sea fluxes for the Arabian Sea yield an annual $N_2O$ flux of 0.33–0.70 Tg $N_2O$. This flux represents approximately 2–35% of the currently estimated global oceanic $N_2O$ source of 2–17 Tg $N_2O$ $yr^{-1}$ (Bange et al., 1996b; Nevison et al., 1995; Suntharalingam and Sarmiento, 2000). The Arabian Sea is the most intensely studied region for $N_2O$ emissions in the world ocean. Given its disproportionately large contribution to this total and the lack of adequate coverage in other potentially important oceanographic regimes, the potential marine contribution to atmospheric $N_2O$ could be somewhat higher than these estimates suggest. Future $N_2O$ flux estimates could be improved by using $N_2O$ concentration data from time series measurements at selected staions in the key regions of the Arabian Sea such as the coastal upwelling areas and the central Arabian Sea.

## Appendix A: Error estimate

In order to evaluate the fit of the computed final $N_2O$ concentrations to the observations, we compared mean annual $1^\circ \times 1^\circ$ data with the smoothed first-guess field and the final field along selected latitudes (Fig. 4). Figure 5 shows the relative error of the predicted seasonal final fields (the so-called interpolation error), estimated as the difference between the final value in each $1^\circ \times 1^\circ$ square and the $1^\circ \times 1^\circ$ pixel data (see Fig. 2). There is a good agreement between predicted values and the observations in the central Arabian Sea during MAM (Fig. 5b). For the monsoon seasons DJF and JJA the relative errors of the predicted values are more variable, indicating a considerable underestimation along the coasts of Oman and southwest India, and an overestimation



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(up to 74%) along the continental shelf of west India (Fig. 5c). The tendency to focus on measurements in locally restricted features such as coastal upwelling in connection with insufficient seasonal data coverage leads to a bias in the first-guess field. For example, the mean for the INDW province, which covers the eastern coastal Arabian Sea, is strongly influenced by high $N_2O$ concentrations observed in the southern Indian continental shelf. In contrast, data coverage for the northern continental shelf is poor, and consequently the first-guess field determines the final $N_2O$ concentration, leading to high uncertainties in this area as indicated by Figs. 5a–5d.

A further uncertainty is introduced by the fact that the $N_2O$ surface concentrations are depending on SST, salinity, atmospheric pressure and the atmospheric $N_2O$ mixing ratio which are, at least partly, subject of long term trends due to global change (Barnett et al., 2001; Levitus et al., 2000). For example, the mean tropospheric $N_2O$ dry mole fractions (see data from the ALE/GAGE/AGAGE program available from the anonymous ftp site cdiac.esd.ornl.edu, subdirectory is given above) increased from about 300 ppb in the late 1970s to about 315 ppb in 1999 suggesting a trend of increasing $N_2O$ surface concentrations. However, a quantification of such trends in sea surface $N_2O$ concentrations is not possible due to the lack of time series measurements in the Arabian Sea. The seasonal northward shift of the Intertropical Convergence Zone introduces air masses of southern hemispheric origin with lower $N_2O$ mole fractions to the Arabian Sea region during the SW monsoon. However, since the mean interhemispheric gradient of $N_2O$ is only about 0.8 ppb (Prather et al., 2001) we did not account for this effect.

## Appendix B: Error propagation

A rough estimate of the mean error of the flux density ($F$), introduced by the uncertainties of the observables (i.e. $T, S, u, P$, and $x'$), was calculated according to the following




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equations:

$$\Delta F = \sqrt{\left(\frac{\partial F}{\partial C_w}\Delta C_w\right)^2 + \left(\frac{\partial F}{\partial C_a}\Delta C_a\right)^2 + \left(\frac{\partial F}{\partial k_w}\Delta k_w\right)^2 + \left(\frac{\partial F}{\partial Sc}\Delta Sc\right)^2}$$

$$\Delta C_a = \sqrt{\left(\frac{\partial C_a}{\partial x'}\Delta x'\right)^2 + \left(\frac{\partial C_a}{\partial \beta}\Delta \beta\right)^2 + \left(\frac{\partial C_a}{\partial P}\Delta P\right)^2}$$

$$\Delta \beta = \sqrt{\left(\frac{\partial \beta}{\partial T}\Delta T\right)^2 + \left(\frac{\partial \beta}{\partial S}\Delta S\right)^2}$$

$$\Delta k_w = \frac{\partial k_w}{\partial u}\Delta u$$

$$\Delta Sc = \sqrt{\left(\frac{\partial Sc}{\partial \boldsymbol{D}_{N2O}}\Delta \boldsymbol{D}_{N2O}\right)^2 + \left(\frac{\partial Sc}{\partial \nu}\Delta \nu\right)^2}$$

where $\nu$ stands for the kinematic viscosity of seawater and the operator $\partial/\partial$ depicts the partial differential. For a strict treatment of the error propagation, the standard deviation of each parameter should be known. Since this was not the case, we replaced the standard deviation partly with best estimates of the mean error (depicted by the $\Delta$ symbol, data listed in Table 5). For $\boldsymbol{\Delta} C_w$ we used the mean relative error (i.e. the interpolation error) calculated from the seasonal data shown in Fig. 5 (see also the Appendix A: Error estimate). We calculated the relative error $\boldsymbol{\Delta} F/F$ for each $1^\circ \times 1^\circ$ square of the four seasonal $N_2O$ fields. Table 6 gives an overview of the resulting mean relative errors of the seasonal flux densities. Not surprisingly, the lowest mean relative error of $C_w$ is associated with highest relative error of resulting flux densities. During MAM the dissolved $N_2O$ concentrations are low and resulting in only small concentration differences ($C_w - C_a$) across the ocean-atmosphere interface which in turn lead to high mean relative errors of the flux densities. During the monsoon season JJA, $N_2O$ concentrations in the coastal upwelling zones are considerable higher causing a higher mean relative error of $C_w$ and comparable low mean relative errors of the resulting flux




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density. The mean relative errors for the seasonal flux densities yield the overall mean relative error of the annual $N_2O$ emissions from the Arabian Sea of at least 65%. Systematic errors caused by uncertainties in parameterizations such as $N_2O$ diffusion in seawater (determination of the $N_2O$ diffusion have not been made in seawater-like systems (see literature compilation in Rhee, 2000)) and air-sea exchange approaches are not accounted for in this estimate (see Results and discussion). Moreover, it is important to keep in mind that the calculation of any climatological data fields are biased by the chosen smoothing and averaging routines (see e.g. Sterl, 2001).

A detailed analysis of errors introduced by different filling routines, averaging procedures etc. is beyond the scope of this study. Generally, gas exchange estimates suffer from the fact that a direct (i.e. at sea) determination of the processes responsible for the gas exchange across the ocean-atmosphere interface is still a technological challenge (Frost and Upstill-Goddard, 1999; Jähne and Haußecker, 1998).

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**Table 1.** Overview of the $N_2O$ Source Data

| Arabian Sea Region | Cruise Dates | Method | N | References |
|---|---|---|---|---|
| West, Central | Dec 1977–Jan 1978 | Con | 668 | Weiss et al. (1992)[a] |
| Northwest, Central | Sep 1986 | Dis | 19 | Law and Owens (1990) |
| East, Central | Dec 1988 | Dis | 15 | Naqvi and Noronha (1991) |
| East, Central | Apr–May 1994, Feb–Mar, Jul–Aug 1995, Aug 1996, Feb 1997 | Dis | 125 | Lal and Patra (1998)[b] |
| Northwest, Central | Sep, Nov–Dec 1994 | Dis | 47 | Upstill-Goddard et al. (1999) |
| Northwest, Central | May, Jul–Aug 1995, Mar, May–Jul 1997 | Con | 1569 | Bange et al. (1996a)[c] Bange et al. (2000)[c] |
| East | Jul 1995 | Dis | 20 | Naqvi et al. (1998) |

Con stands for continuous measurements.
Dis stands for measurements of discrete samples. N stands for number of data points.
[a] Data are available from the anonymous ftp site cdiac.esd.ornl.edu (subdirectory /pub/ndp044) at the Carbon Dioxide Information Analysis Center in Oak Ridge, Tennessee.
[b] Data are included in the JGOFS-India data compilation on CD-ROM available from the Indian National Oceanographic Data Centre, Goa, India (ocean@csnio.ren.nic.in).
[c] Data are available from the German JGOFS data management (http://www.ifm.uni-kiel.de/jgofs/dm).




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**Table 2.** $N_2O$ fluxes from the Arabian Sea calculated with the $N_2O$ diffusion coefficient of Broecker and Peng (1974)

| $N_2O$ Fields | Flux,[a] Tg $N_2O$ | Percentage,[a] % |
|---|---|---|
| DJF | 0.08 / 0.13 / 0.19 | 22 / 25 / 24 |
| MAM | 0.01 / 0.01 / 0.02 | 3 / 2 / 3 |
| JJA | 0.25 / 0.33 / 0.51 | 68 / 65 / 65 |
| SON | 0.03 / 0.04 / 0.06 | 8 / 8 / 8 |
| Sum | 0.37 / 0.51 / 0.78 | |

[a] First value calculated according to LM86; second value calculated according to N00, and third value calculated according to W92.




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**Table 3.** $N_2O$ Fluxes from the Arabian Sea calculated with the $N_2O$ diffusion coefficient of Rhee (2000)

| $N_2O$ Fields | Flux,[a] Tg $N_2O$ | Percentage,[a] % |
|---|---|---|
| DJF | 0.07 / 0.12 / 0.17 | 21 / 26 / 24 |
| MAM | 0.01 / 0.01 / 0.02 | 3 / 2 / 3 |
| JJA | 0.23 / 0.30 / 0.45 | 70 / 64 / 64 |
| SON | 0.02 / 0.04 / 0.06 | 6 / 9 / 9 |
| Sum | 0.33 / 0.47 / 0.70 | |

[a] First value calculated according to LM86; second value calculated according to N00, and third value calculated according to W92.



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**Table 4.** Summary of various $N_2O$ flux estimates for the Arabian Sea

| Source region | Area, $10^6$ km$^2$ | Flux, Tg $N_2O$ yr$^{-1}$ | References |
|---|---|---|---|
| Central, west (>15° N) | 1.6 | 0.22–0.39 | Law and Owens (1990) |
| Central, east | 6.2 | 0.44 | Naqvi and Noronha (1991) |
| Central, west | 6.2 | 0.8–1.5 | Bange et al. (1996a) |
| Central, east | 6.2 | 0.56–1.00 | Lal and Patra (1998) |
| Central, west | 8.0 | (0.41–0.75)[a] | Upstill-Goddard et al. (1999) |
| Central, west (> 6° N) | 4.9 | 0.16–0.31 | Bange et al. (2000) |
| >15° N | 1.6 | 0.10–0.21 | This study[b] |
| > 6° N | 4.4 | 0.28–0.60 | |
| > Equator | 6.8 | 0.37[c]–0.78[c] | |

[a] Semi-annual flux.
[b] Data calculated with the diffusion coefficient of Broecker and Peng (1974). First value calculated according to LM86; second value calculated according to W92.
[c] Taken from Table 2.




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**Table 5.** Errors used for the error propagation

| Observable quantity | Mean error | References, remarks |
|---|---|---|
| Water temperature, $T$ | $\pm$ 0.5 K | McClain et al. (1985) |
| Salinity, $S$ | $\pm$ 0.1 | Estimate |
| $N_2O$ dry mole fraction, $x'$ | $\pm$ 2% | Estimate |
| Atmospheric pressure, $P$ | $\pm$ 5% | Estimate |
| Kinematic viscosity, $\nu$ | $\pm$ 1% | Estimate |
| Diffusion of $N_2O$, $D$ | $\pm$ 10% | Estimate |
| Wind speed, $u$ | $\pm$ 1.4 m $s^{-1}$ | Schlüssel (1995) |



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**Table 6.** Overview of the mean relative errors of $N_2O$ surface concentrations and $N_2O$ flux densities

| | Mean $\Delta C_w/C_w$, $\pm$ % | Mean $\Delta F/F$, $\pm$ % |
|---|---|---|
| DJF | 11 | 75 |
| MAM | 4 | 330 |
| JJA | 14 | 79 |
| SON | 12 | 442 |




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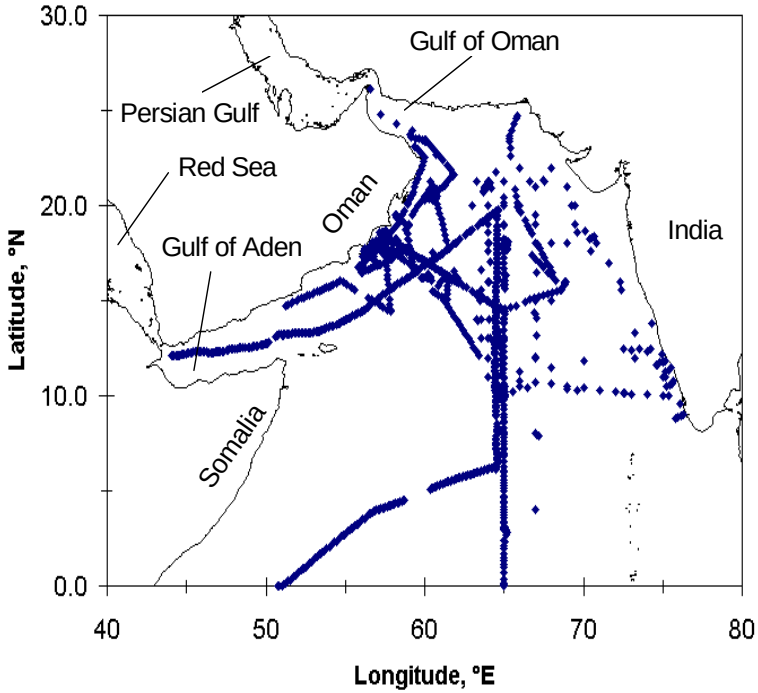


**Fig. 1.** Map of the Arabian Sea with locations of the $N_2O$ measurements in the surface layer used in our study (see Table 1).



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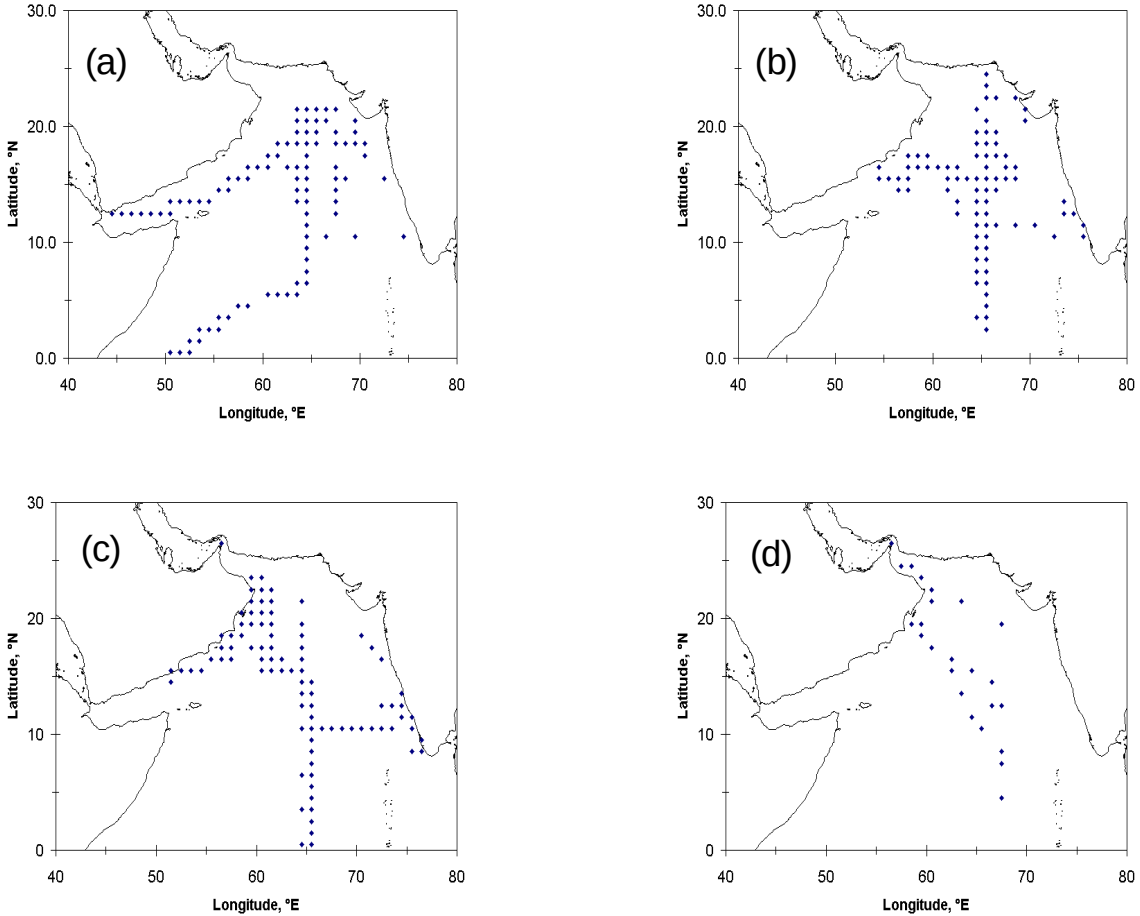


**Fig. 2.** Seasonal maps of $N_2O$ pixels. **(a)** DJF, **(b)** MAM, **(c)** JJA, and **(d)** SON.



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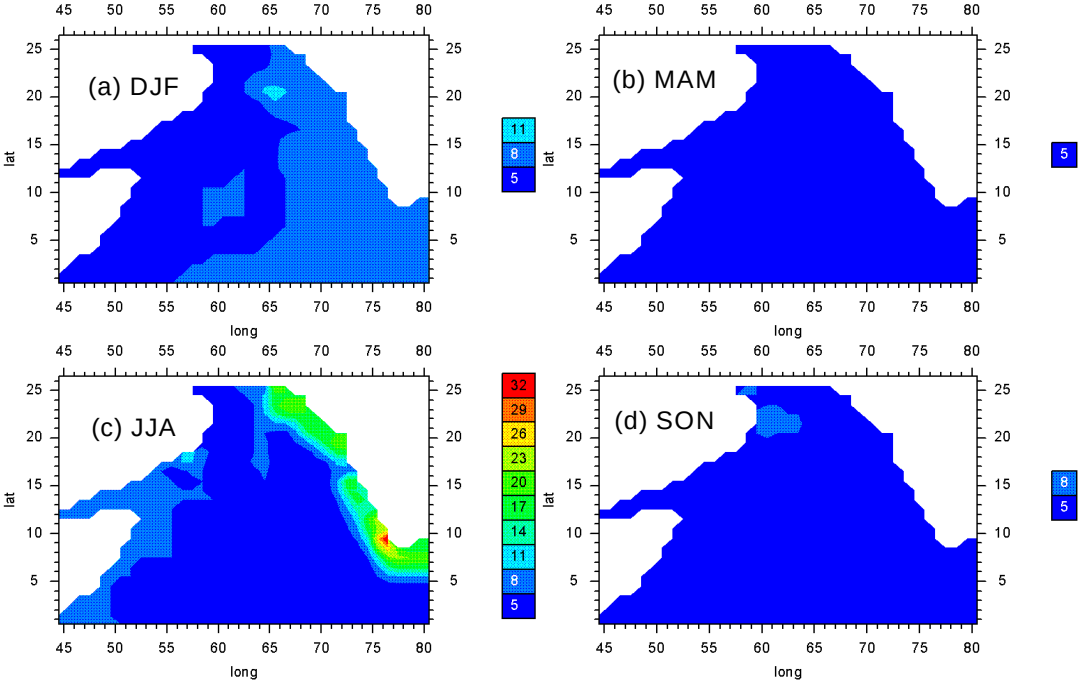


**Fig. 3.** Seasonal maps of the final $N_2O$ $1^\circ \times 1^\circ$ fields (in nmol $L^{-1}$). **(a)** DJF, **(b)** MAM, **(c)** JJA, and **(d)** SON. Contour labelling starts with 5 nmol $L^{-1}$; minimum concentration range is shown in dark blue (5-8 nmol $L^{-1}$), maximum concentration range is shown in red ($>$ 32 nmol $L^{-1}$).



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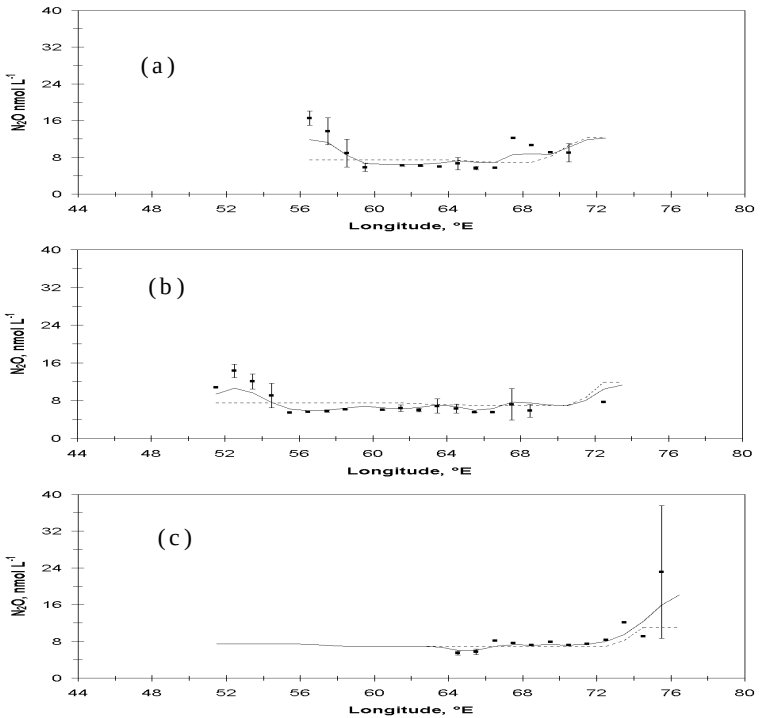


**Fig. 4.** Annual mean $N_2O$ concentrations (in nmol $L^{-1}$) along selected latitudes. The solid line is the predicted $N_2O$ from the final $1^\circ \times 1^\circ$ field, the dashed line stands for the smoothed first-guess field, and the solid squares represent the annual mean $N_2O$ with standard deviation of all measurements within the $1^\circ \times 1^\circ$ squares along the given latitude. (When less than 3 values were available no standard deviation is given.) **(a)** 18.5° N, **(b)** 15.5° N, and **(c)** 10.5° N.




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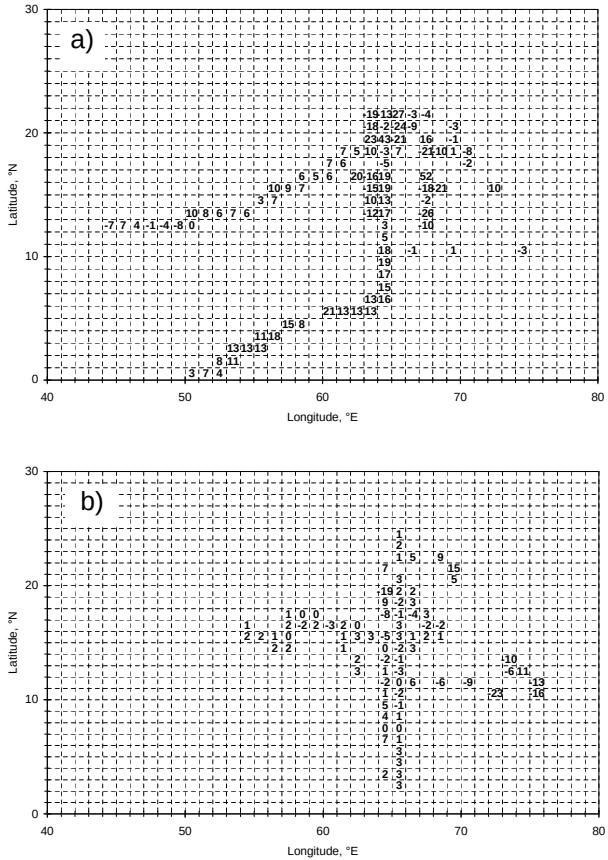


**Fig. 5.** Relative errors of the final field 1°× 1° values. **(a)** DJF, **(b)** MAM, **(c)** JJA, and **(d)** SON.



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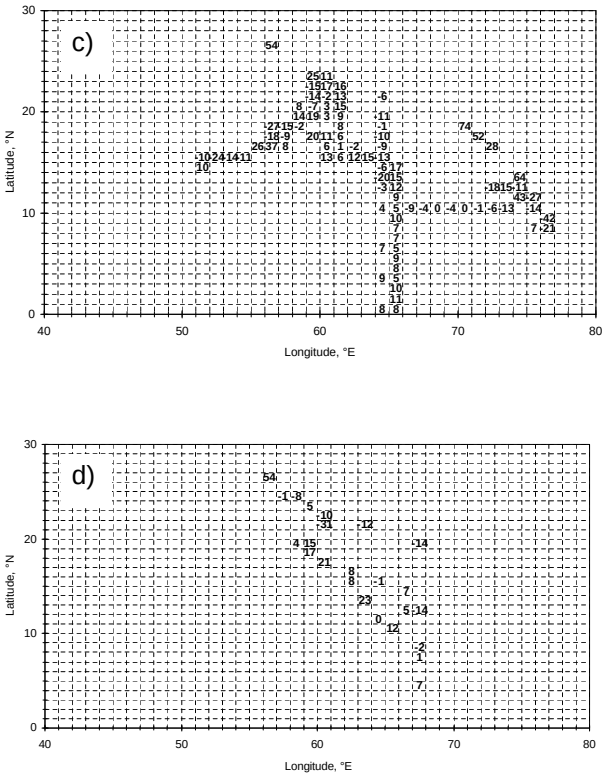


**Fig. 5. Continued.** Relative errors of the final field 1° × 1° values. **(a)** DJF, **(b)** MAM, **(c)** JJA, and **(d)** SON.



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